AN INTEGRATED EXPLORATION APPROACH TO MAP BIF-HOSTED GOLD DEPOSITS IN THE RIO DAS VELHAS GREENSTONE BELT, QUADRILÁTERO FERRÍFERO, BRAZIL

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INTRODUCTION

In the Quadrilátero Ferrífero (QF) region, many exploration surveys have been conducted based on old prospects, geological and geochemical mapping. These efforts were the primary tools responsible for many gold discoveries. Limited outcrops and a thick soils pose serious logistical constraints to finding new deposits in this region (Ladeira 1980, Pinto 1996, Silva 1999). Because of the extensive cover, there is a higher than usual likelihood that undiscovered deposits exists. Hidden mineral deposits are prime targets for geophysical methods, which image the third dimension that mapping geologists can only infer. The availability of airborne geophysical survey data combined with GIS (geographic information system) now allow emphasis to be placed on mapping as well as subsurface features related to mineralization in the Quadrilátero Ferrífero region (Silva 1999). Most mineral deposits are spatially small part (less than 3km²) of a mineralized system (Jaques et al. 1997), which can occur over a scale of tens of kilometers (district scale) or hundreds of kilometers (regional scale). The Rio das Velhas Greenstone Belt is located within the QF and is characterized by world-class gold production in regard to deposits hosted in banded iron formation of the Algoma type. In the Rio das Velhas Greenstone Belt case, much of the exploration effort has focused on deposits with obvious surface expression. Such deposits commonly have a regional and/or local magnetic and/or radiometric and/or electromagnetic signature, many of which were mapped by Rio das Velhas airborne survey.

Although spatial data integration methods are being used increasingly for mineral resource assessments, a preferred methodology is by no means established. The purpose of combining various geologic, geophysical, and geochemical information is to identify and describe spatial associations present in the data, and to develop models for analysis and prediction of spatial phenom-
This paper discusses a quantitative data integration approach for establishing areas of mineral potential for oxide-type BIF of Rio das Velhas Greenstone Belt, Quadrilátero Ferrífero, currently an important area for gold exploration. The general goal of this study is to combine information from the known geology, airborne geophysics, and mineral occurrences to map areas favorable for gold mineralization on an intermediate regional scale (1:100,000). Emphasis is placed on predicting the distribution of BIF-Archean-hosted-gold mineralization using the probability ratio method (Silva 1999, Lee et al. 2001). Results of the predictive models are evaluated with ground geologic field investigations to establish their validity.

**GEOLOGIC OVERVIEW** The QF is located in the southern portion of the São Francisco Craton and is composed of Archean granite-gneiss terrains (GGTs), Archean Greenstone Belt (Rio das Velhas Supergroup), Paleoproterozoic (Minas Supergroup and Itacolomi Group) and Paleoproterozoic-Mesoproterozoic (Espinhaço Supergroup) supracrustal units (Fig. 1).

The supracrustal units of the Minas and Rio das Velhas Supergroups, surround and are surrounded by granite-gneiss domes and each has been named: Bação, Caeté, Bonfim, Belo Horizonte and Santa Rita. These domes consist of poly-deformed gneiss, metatonalites to metagranites, amphibolites, meta-ultramafic rocks, as well as pegmatites formed under amphibolite facies conditions during the Archean and the Transamazonic Eras (Cordani et al. 1980). The contact with adjacent supracrustal units is tectonic.

The Rio das Velhas Supergroup (RVSG – 3.0-2.7 Ga) is divided in two groups. The first, Nova Lima Group (NGL) comprises a lower ultramafic unit, an intermediate felsic-mafic unit and a clastic-mafic-felsic unit (Ladeira 1980). These rocks are overlain by quartzites of Maquiné Group.

Zuchetti et al. (1996) divided the Rio das Velhas Supergroup into informal lithostratigraphic units based on field observations, petrologic and geochemical studies. Pinto & Silva (1996) proposed the upgrading of these units to formation status. The units were grouped into four following tectonic blocks (Pinto 1996): Nova Lima, Caeté, Santa Bárbara, and São Bartolomeu. These blocks represent distinct petrogenetic environments.

Along with these modifications, the formal designation of the Rio das Velhas Supergroup was kept, with the Nova Lima Group (bottom) and Maquiné (top). On the other hand, the Córrego dos Boiadeiros Complex and the Quebra Osso Group were redefined as formations within the Nova Lima Group. The Córrego dos Boiadeiros Complex was separated as an intrusive ultrabasic sill at the bottom of the Nova Lima Group (Pinto 1996, Zuchetti et al. 1996, Pinto & Silva 1996). According to Pinto and Silva (1996), three formations (Córrego Ouro Fino, Córrego dos Boiadeiros, and Quebra Osso Formations) belong at the base of the Nova Lima Group and are present in the Caeté, Nova Lima, and Santa Bárbara Blocks (Fig. 2).

Volcanogenic metasedimentary rocks of the Nova Lima Group are only exposed in the Caeté and Nova Lima Blocks and are represented by the intercalation of felsic pyroclastic, and epiclastic horizons. Gradational layering, small cross stratification and turbiditic structures are commonly preserved. Fine-grained carbonaceous schists also occur and are interpreted as metapelites. They are intercalated in the sequence and were designated as the Lapa Seca Formation by Ladeira (1980). On the other hand, the volcanoclastic metasedimentary rocks occur in the Caeté Block and were named the Ribeirão Vermelho Formation. This formation consists of pyroclastic dacitic tuff and agglomerate with subordinated flows flows (Pinto & Silva 1996, Fig. 2).

The metasedimentary rocks were previously assigned to six formations: Córrego do Sítio (Caeté, Nova Lima and Santa Bárbara Blocks), Mindá (Caeté, Nova Lima and Santa Bárbara Blocks), Catarina Mendes, Córrego da Paina, Fazenda Velha and Pau D’Óleo (São Bartolomeu Block) (Pinto & Silva 1996, Fig. 2).

**Tectonic Evolution** The Quadrilátero Ferrífero is marked by a complex multi-phased history, which resulted in the heterogeneous superposition of successive tectonic events. Granite-gneiss terrains older than 3000 Ma are considerably more complex than younger ones. In addition, younger ones that have yielded 2700 to 2800 Ma ages have a NS foliation with horizontal mineral lineation, which formed coevally with the NS structures of QF, which are well displayed in the Bonfim Complex (Chemale Jr. et al. 1991, 1994).

Endo & Carneiro (1996) proposed that the deformation and its successive reactivations, during the Neorarchean, was formed within a transpressional tectonic regime. The NS structures are deflected in some places, such as in the Bação Complex and surrounding features (Endo et al. 1996a). Such deflection was interpreted by Endo et al. (1996a) as the result of the Bação Tectonic Complex presence that would have nucleated during the crustal
fragmentation episodes. In this context, the mylonitic foliation on the margins of the complex, along with the dextral sigmoidal trajectory of the mylonitic foliation within the Rio das Velhas Super-
group, would have been generated during the oldest dextral de-
formation episode, about 2780 Ma.

Marshak et al. (1992) proposed that the Rio das Velhas Green-
stone Belt possesses dome-and-keel structures. The supracrustal
sequences having been placed in narrow synformal depressions
between domal metamorphic complexes. These authors demon-
strated that the dome-and-keel structures are younger than the
volcano- sedimentary sequence, and the majority of the younger
GGT’s granitoids. The structures involved the Paleoproterozoic
Minas Supergroup. According to Marshak et al. (1992), the gen-
eration of dome-and-keel structures would occur through the al-
most-diapiric ascension of the metamorphic complex bodies in an
extensional environment associated with partial melting that gen-
erated felsic magmas, which in turn formed contact metamorphic
aureoles in the intruded supracrustal rocks. Chemale Jr. et al. (1991
and 1994) interpreted the domes of the Quadrilátero Ferrífero to
have formed about 2.0 Ga as the result of crustal extension, similar
to the formation of the metamorphic core complexes in the south-
western United States.

Gold Mineralization According to Ribeiro-Rodrigues (1998),
the QF gold production reached approximately 910t. Archean green-
stone-hosted deposits are responsible for 53% of this production
and constitute the majority of the inactive and active mines, such as:
Cuiabá (Ribeiro-Rodrigues et al. 1996 a, b, c and Ribeiro
Rodrigues 1998), São Bento, Raposos (Godoy 1994), Lamego
(Carmo in prep.), Morro Velho (Ladeira 1988), Côrrego do Sítio
(Takai et al. 1991) and Juca Vieira. Proterozoic sediment-hosted
deposits have contributed approximately 8% of the gold produc-
tion, while Cenozoic deposits have contributed almost 39% of the
production.

Banded iron formations are by far the most important hosts of
gold mineralization. They can be divided into three types: oxide
type (Raposos, Espírito Santo), carbonate-sulphide type (Cuiabá,
Lamego) and ankerite/ferroan dolomite, quartz and plagioclase type
(Lapa Seca – Morro Velho, Bicalho, Juca Vieira). These three types
account for 99% of the gold production (approximately 540 t).
They are associated mainly mafic volcanics (metabasalts and meta-
andesites) and carbonated pelites, which, sometimes are mineralized
adjacent to banded iron formations.

In terms of the amount of contained gold, the Lapa Seca (Mestre
Caetano Unit) is the major host. Also, BIF and chert host significant
economic grade-mineralization and account for 14% of the total
gold production (Ribeiro-Rodrigues 1998). The common wallrock
alteration is sulphidation and silicification in the BIF’s, sericitization,
carbonatization, chloritization, and sulphidation in the mafic rocks,
and carbonatization, sericitization, and sulphidation in the
metasedimentary rocks. However, in terms of the number of indivi-
dual deposits, BIF is the principal host, with 62% of the deposits.

STATISTICAL METHODOLOGY AND DATA A method based
on comparative statistics is applied to grids of the airborne geo-
physical data to determine whether or not there exists a character-
istic signature over particular geologic units with known high
potential for gold mineralization. In general, the method calculates
a ratio of probabilities, defined here as weights, which numerically describe how well a particular evidential data layer predicts a given training area. The higher in value the weight, the stronger the spatial association and the more likely the particular evidential layer predicts the location of a given training region.

THE DATA SETS A variety of regional data sets from Rio das Velhas Greenstone Belt have been registered and analyzed using a geographic information system (GIS). The data sets include Rio das Velhas Project geological map (reference), Rio das Velhas airborne geophysical survey data (airborne magnetic, radiometric and frequency domain electromagnetics), and terrain. Gridded aeromagnetic, electromagnetic and four channel radiometrics data were used as evidential layers to predict favourable host rocks and mineralized environments (Silva 1999 and Silva et al., in this issue). As described in Silva (1999) and Silva et al. (this issue), different steps were employed to enhance the predictive layers and prepare them for modeling. The data used as evidential layers and the choice of training areas were selected to reflect as much as possible some of the current ideas about gold genesis in the area.

The calculation of probability ratios was carried out on raster versions of the geologic, mine, geophysical data represented on 50-meter pixels with ER Mapper 6.0 and ERDAS Imagine 8.3. Different classes of each information layer were considered candidates for this analysis. Some assumptions are considered in these cases:

- The geological map of the study area and the host rocks (Morro Vermelho and Santa Quitéria units) are known and used as predictor of mineral potential of oxide BIF-hosted-gold mineralization. The goal is to define the geophysical signature around host rock and then extend this information into unknown areas.
- The resulting weights represents the measures of the spatial correlation between the host rock and input data set (geophysical data). Values greater than 1 indicate a positive association between the tested class and tested domain, a value less than 1 suggests a negative association and a value equal one implies a random association. For example, a weight of 1.2 for a class implies that class is 12 times more likely to be found within the reference area as in the other (non-reference) areas in the study area.

Calculation of probability ratios Probability ratios were determined for each input map and reflect the correlation between host rocks with occurrences and deposits. After probability ratios were calculated, a composite model was created showing the probability of occurrence based on accumulated evidence from the input maps. In simplified terms, probability ratios can also be represented as follows:

\[
PR = \frac{\text{Overlap area}}{\text{Tested Class Area}}
\]

Prototype area/Study Area

Where the overlap area is defined as the overlap between the prototype (favourable geologic formations) and the tested class (classes in the geophysical data). The prototype area for this study is the oxide-type banded iron formation (BIF) or favourable geologic-host units with great amount of oxide-type BIF. The study area is the area covered by the airborne geophysical survey. The expression above is equivalent to the ratio of two probabilities. The probability ratios can be expressed as the following formula:

\[
PR = \frac{\text{Conditional Probability}}{\text{Prior Probability}}
\]

The conditional or posterior probability (numerator) represents the odds of finding the test class within a reference area. The prior probability or unconditional probability (denominator) describes the odds of being in the prototype area if no conditions are known about the location of the tested class, except that the prototype is located somewhere within the study area. We measure the relative likelihood of finding the prototype within each tested class area compared to the likelihood of finding the prototype within the study area (McCafferty et al. 1999, Silva 1999).

One major assumption is made prior to calculating additive predictive models: that the input maps are conditionally independent. This assumption is important for models in which weights of different layers are to be combined additively. If conditional dependence is present among two or more layers, the resulting model values will be inflated by essentially double weighting of conditionally dependent variables. If the expert uses input maps that are not related, the conditional dependence is automatically avoided. Furthermore, although it is probably unlikely that conditional dependence can be completely, it is not devastating to the results of modeling if the expert is using the models to indicate favorability or ranked potential. Bonham-Carter et al. (1989) suggest that this assumption can be tested. Either chi-square or pairwise tests can be applied, showing the degree with which each input maps satisfies the conditional independence with respect to the prototype areas.

DEVELOPMENT OF PREDICTIVE MODELS FOR BIF-HOSTED GOLD Geologic units host large deposits of oxide-BIF gold mineralization, which include the Morro Vermelho and Santa Quitéria units (Silva 1999, Fig. 2), are used as training regions. We use the airborne geophysical data (magnetic, electromagnetic, and radiometric) as the evidential layers (Silva 1999).

Our initial hypothesis is that BIF units will likely have a characteristic geophysical signature that can be used to predict the locations of similar types of lithologies. Because of the ability of the magnetic and electromagnetic data to image through vegetation and soil cover, we expect that many of the areas mapped by the approach may locate similar types of BIF at depth.

Fundamental lithological, mineralogical, and structural controls on gold mineralization are summarized in order to establish a geologic framework from which the geophysical predictive models were built. This section describes the ”expert” knowledge component of the model building efforts and results primarily from extensive literature research and secondary field studies of banded iron-hosted gold deposits in the Rio das Velhas study area.

Lithological Controls The Archean greenstone-hosted gold deposits are related to the volcano-sedimentary successions of the Nova Lima Group, mainly with Fe-rich lithologies such as carbonate-facies iron-formations, and mafic volcanics. Those deposits occur at the ultramafic, felsic and mafic volcanics, clastic and chemical sediments and are characterized by common features (Ribeiro-Rodrigues 1998).

The common wallrock alteration is the sulphidation and silicification in the BIFs, sericitization, carbonatization, chloritization and sulphidation in the mafic rocks and carbonatization, sericitization and sulphidation in the metassediments.

The oxide-banded iron formation was studied, because it is an important host and occurs in different positions within the Nova Lima Group. The objective in this case is try to map the likelihood of BIF inside of Rio das Velhas Greenstone Belt.

The banded iron formations are by far the most important hosts
of economically viable gold deposits and have been contributing 99% of the whole gold production. These banded iron formations have different geological and different physical properties and can be divided in three types: oxide, carbonate-sulfide and ankerite/ferroan dolomite, quartz and plagioclase type (Lapa Seca). In terms of the amount of contained gold, the Lapa Seca (interlayered with the Mestre Caetano Formation) is the major host lithology. Oxide-type BIFs, containing significant economic grade-mineralization, host 62% of the known gold-bearing deposits in the area.

Mineralogical Controls Usually, the ore is composed of pyrite, pyrrhotite, arsenopyrite, sometimes, with smaller quantities of calcocite, sfarelite and galena. The percentages of these mineral phases vary not only among different deposits, but also within the deposits themselves. Gold is associated with sulfides, occurring as inclusions, in fractures or along grain boundaries (Ribeiro Rodrigues 1998).

Structural Controls The Quadrilátero Ferrífero has a complex multi-phase history characterized by the heterogenous superposition of successive tectonic events, which is an important factor in controlling the distribution of the gold. One important feature that indicates structural control is the remarkable down-plunge continuity of the ore bodies, parallel to the stretching lineation, and mineralization in shear zones within metavolcanics (Veira 1991a and 1991b, Scarpelli 1991, Ribeiro-Rodrigues et al. 1996 a, b, c, Ribeiro-Rodrigues 1998).

PREDICTIVE MODELS FOR FAVOURABLE HOST ROCKS

Weights that describe the degree or strength of the spatial association between various geophysical data layers with the geologic units of the Morro Vermelho and Santa Quitéria Formations were calculated using the procedure given in Silva (1999). Integrated predictive models that combine a number of statistically significant geophysical data are presented with a description of the mathematical and geologic rationale.

Predictive Model Construction Probability ratios were calculated to determine if the geologic units with known potential for gold mineralization have characteristic magnetic, electromagnetic and radiometric signatures. An initial effort was made to calculate probability ratios for all geophysical and derivative data layers. Analyses were conducted among RTP, derivative products of RTP, analytic signal amplitude, K, U, Th, total count and conductivity data and their derivative products.

Although probability ratios were calculated for all the geophysical data layers, only some products were determined to have spatial association with the select geologic units and incorporated in final models.

Predictive models for favorable host rocks and mineralized environments were generated for this study using the sum-of-weights approach. This approach was successfully used to generate predictive models for a regional study in Montana (Lee et al. 2001) and a mineral resource study in southeast Alaska (McCafferty et al. 1999). Co-registered pixels were added to create composites of multiple layers of geophysical and radiometric information. Only those pixels with probability ratios greater than 2 were considered for incorporation into final models. Additionally, each pixel present in final models that incorporated more than one evidential layer represents a either sum of probability ratios between multiple data layers or null values. For this reason, the minimum weight for any predictive model is 4 (2 + 2). For example, for pixels where there is a probability ratio for conductivity but not for magnetization, these pixels are assigned a null value. In the summing process, one of the weights is rejected and is assigned a null value in final predictive model.

**BIF–HOSTED GOLD PREDICTIVE MODEL** Oxide-banded-iron formation, Morro Vermelho Formation, and Santa Quitéria Formation predictive geophysical model illustrate one of the best models for the study area (Fig. 2). These units are represented in both the magnetization (terrace of high pass) and conductivity maps as high values with relatively high probability ratios indicating strong spatial associations with the oxide banded-iron-formation (Fig. 3) (Silva 1999 and Silva et al., this issue).

The terracing technique was applied to aeromagnetic data after it was reduced-to-pole to create a terrace-magnetization map. However, the long wavelengths, present in the reduced-to-pole data, did not permit to see the shorter wavelength anomalies corresponding to important features related to mineralization. Because of that, the terrace filter was applied to the high-pass. The goal was to point out the form and amplitude of magnetic responses of mineral deposits in the Rio das Velhas Greenstone Belt, where the highest concentration of the important mineralized sites is located.

Grid of 4175 Hz resistivity converted to conductivity was terraced to produce inferred conductivity maps. These maps emphasize abrupt boundaries between rocks with different conductivities. Map emphasizes boundaries between different rocks with varying conductivities. The most important host rocks or mineralized domains show high conductivity response. There is a strong correlation of these domains with the mapped Morro Vermelho and Santa Quitéria Formations. Such domains that host the known gold mineralization, illustrating the utility of this layer to improve the model and the understanding of mineralization.

Analysis of the radiometric data indicated no spatial association with these geologic units. Probability ratios are all less than 2 for potassium, uranium, and thorium indicating no characteristic radiometric signature. The uranium shows depletion in the high concentration range values, which is characteristic of tropical terrains. Since none of the three radiometric channels had weights greater than 2 these data layers were not incorporated in the final predictive model.

Figure 3 shows a predictive model for oxide-BIF geologic units based on the analysis of derivative products calculated from magnetics and 4175Hz-electromagnetics. A sum-of-weights model was created from the magnetization and conductivity data with weights that range from 4 to 25. Areas mapped on Figure 3 represent tracts of ground with favourable conditions to map oxide BIF-host rocks. The modeled areas similar geophysical properties as the known oxide BIF geologic units.

**Field Verification of Predictive Model for Oxide-Type Banded Iron Formation** A modeled area of oxide-BIF was ground truthed with geologic fieldwork (Fig. 4). Fieldwork discovered an unmapped exposure of banded iron formation hosted in the Córrego do Sítio Formation. The geologic exposure is smaller than the model predicts. Most likely, the unit extends into the shallow subsurface and demonstrates the ability of the probability ratio technique to map shallow subsurface geology. Although ground follow-up and drill data is required to verify this inference, the approach shows encouraging potential to identify possible shallowly buried favourable host.
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Figure 3 - Predictive geophysical model for hosts of magnetic banded iron gold mineralization. Probability ratios range from 4 (blue) to 25 (red) and would define combinations of high magnetizations and conductivities values that are 4 to 25 times more likely to map the magnetic BIF than any other geologic unit in airborne survey area. Light gray shows mapped extent of geologic units with strong associations with magnetic BIF hosted gold mineralization. Select mines names are labeled, as is the potential target for host of gold mineralization.

Figure 4 - Oxide-banded iron formation hosted in the Córrego do Sítio Formation, northwest of the Rio das Velhas Greenstone Belt. Target mapped with the predictive geophysical model presented in the Fig. 3.
This model illustrates an association in the geophysical data that helps characterize geological units important to mineralization (Fig. 3). Through the probability ratio mapping approach, it is possible to observe the geologic units that have strong spatial associations with oxide-banded iron formation, chemical exhalative and volcaniclastic metasediments.

**CONCLUSIONS AND TARGET CONCEPTIONS** This paper focuses mainly on development of the probability ratio mapping technique and its application on prediction of the distribution of oxide-type BIF-hosted gold deposits in the Rio das Velhas Greenstone Belt, Quadrilátero Ferrífero. The probability ratio models use the spatial distribution of host rocks to calculate different multi-map signatures for gold mineralization.

The ability of the models to predict regions favorable to gold mineralization was verified by comparing them with the occurrences of Archean-greenstone-hosted gold deposits and host rocks (Silva, 1999, Silva et al. 1999). From the regional perspective, the predictive geophysical model convincingly defines the known host rocks. Most importantly, the predictive model delineate prospective areas different from those in which are the known mines. These areas have the potential to contain similar mineralization.

The predictive model leads to a consideration about the future mineral exploration efforts in the Rio das Velhas Greenstone Belt. The characterization of favorable host rocks and several of the large current and past-producing gold mines in the area provide tools for this discussion. If the veracity of these predictive models is ground truthed and confirmed, then the results will help to conduct new exploration activities in the region. This approach should consider high-potential host rocks, such as extension areas of highly prospective nature and potential targets mapped by this technique.

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**References**


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